

# **Nigerian Journal of Soil Science**

Journal homepage:www.soilsjournalnigeria.com



# ASSESSMENT OF DITHIONITE AND OXALATE EXTRACTABLE IRON AND ALUMINIUM OXIDES IN SOILS SUPPORTING RAPHIA PALMS AT NIFOR MAIN STATION

\*Osayande P.E., Oviasogie P.O., Aisueni N.O., Stephen O., Irhemu P. and Ekebafe M.O. Chemistry Division, Nigerian Institute for Oil Palm Research, P.M.B. 1030, Benin City-Nigeria \*Corresponding author's email: <a href="mailto:pvefosa2001@yahoo.co.uk">pvefosa2001@yahoo.co.uk</a>,

# **ABSTRACT:**

The oxides of iron and aluminium are known to affect the physical and chemical properties of soils such as ion adsorption, aggregate stability, phosphorus sorption and charge characteristics. This study was conducted to provide information on the status of free oxides in both soils of the Inland Valley and a well drained Upland soil at the NIFOR main station. Twenty four (24) representative samples were obtained from twelve (12) mini pits sunk at the two locations. The samples were air-dried and analyzed for physico-chemical properties including iron and aluminium oxides using standard methods. The results indicated that Fe<sub>d</sub> ranged from 5.86mg/kg to 151.38mg/kg at the Inland Valley and ranged from 11.13mg/kg to 34.46mg/kg at the Upland. The Fe<sub>d</sub> values were higher than Fe<sub>o</sub> values in both locations. The Fe<sub>o</sub>/Fe<sub>d</sub> ratio (diagnostic index) values clearly distinguished the well drained Upland soil from the poorly drained Inland Valley as the values decreased with increased soil depth at the Inland Valley but increased significantly with increased soil depth at the Upland. There was no evidence of co-migration of clay with Fed in both locations though significant correlations were obtained between clay and clay/Fe<sub>d</sub> ratio of Inland Valley and Upland with (r = 0.841, P < 0.01) and (r = 0.891, P < 0.01) respectively. Significant correlations were also obtained between Fe<sub>d</sub> and Fe<sub>d</sub>-Fe<sub>o</sub> (r = 0.998, P < 0.01); Fe<sub>d</sub> and organic carbon (r = 0.945, P < 0.01); Fe<sub>d</sub> and clay (r = -0.720, P < 0.01); Fe<sub>d</sub> and Clay/Fe<sub>d</sub> ratio (r = -0.719, P < 0.01) Al<sub>d</sub> and Al<sub>d</sub>-Al<sub>o</sub> (r = 0.922, P < 0.01); Al<sub>d</sub> and Fe<sub>o</sub>/Fe<sub>d</sub> (r = 0.592, P <0.05); Al<sub>d</sub> and Clay/Fe<sub>d</sub> (r = 0.628, P< 0.05); Al<sub>o</sub> and Al<sub>d</sub>-Al<sub>o</sub> (r = -0.653, P < 0.05); Fe<sub>d</sub>-Fe<sub>o</sub> and Clay/Fe<sub>d</sub>(r = -0.724, P < 0.01); Fe<sub>d</sub>-Fe<sub>o</sub> and organic carbon (r = 0.948, P < 0.01); Fe<sub>d</sub>-Fe<sub>o</sub> and Clay (r = -0.736, P 0.01); Fe<sub>o</sub>/Fe<sub>d</sub> and Clay/Fe<sub>d</sub> (r = 0.617, P < 0.05) and Fe<sub>o</sub>/Fe<sub>d</sub> and Clay ( r = 0.705, P < 0.01) at the Inland Valley. This trend of correlation continued at the Upland except that Al<sub>d</sub> was not correlated with Fe<sub>o</sub>/Fe<sub>d</sub> and Clay/Fe<sub>d</sub> while Fe<sub>o</sub> was correlated with Fe<sub>o</sub>/Fe<sub>d</sub> of Upland (r = 0.607, P < 0.05) and was not correlated with Fe<sub>0</sub>/Fe<sub>d</sub> of the Inland Valley. This study has revealed that the soils of NIFOR main station are highly weathered soils and contained little weatherable minerals as revealed by the low amounts of Fe<sub>0</sub>. The low amounts of the crystalline oxides in these soils reveal a lack of accumulation of concretions, nodules and plinthites which lead to deterioration of physical and chemical properties of soils.

Key words: Dithionite, NIFOR, Oxalate, Raphia palms, Sesquioxides

#### INTRODUCTION

Many soils in Nigeria have been reported to contain sesquioxides in varying proportions (Ogunkunle and Onansaya, 1992; Agbenin, The crystalline forms of the 2003). sesquioxides had been found to dominate in the basement complex soils of Southwestern Nigeria (Ogunkunle and Onansaya, 1992). This is the more advanced stage compared to the amorphous forms that are mobile in the soil and could be associated with organic matter (Obi et al., 2009). Percentages of free iron (Fe<sup>3+</sup>) have long been used as aids in distinguishing soil types and differentiating soil horizons (McKeague et al., 1971). Comparisons of percentages of free iron frequently have been used as aids in determining relative ages of soils (Ogunsola et al., 1989). Iron is subject to increase in concentration through soil development processes, as reflected by contents of ten to fifteen percent in many soil colloids (Osodeke et al., 2005). Soils having the high percentages are approaching the iron content of low grade iron ore. High stability of soil aggregates and high soil porosity are usually associated with high iron oxide contents. Two analytical methods are commonly used in estimating percentages of free iron in soils: the acidammonium oxalate method first applied to soils by Schwertmann (1964) and the dithionite-citrate-bicarbonate method proposed by Mehra and Jackson (1960). The oxalate extraction is presumed to remove the poorly dithionite crystalline iron oxides. The extraction is presumed to remove finely crystalline iron oxides in addition to the oxalate-soluble fractions. Therefore, amount of iron released by the dithionite method (Fe<sub>d</sub>) should be equal to or greater than the amount of iron released by the oxalate method (Fe<sub>0</sub>). The difference between the values obtained by the two methods represents the amount of iron present in definite crystalline forms (Ogunsola et al., 1989) so that with increasing soil age, the crystalline iron oxides increase at the expense of the poorly crystalline forms. This trend is reflected

in the ratio of Fe<sub>o</sub> to Fe<sub>d</sub> (Fe ratio) (Ogunsola et al. 1989). Ideally, the Fe ratio has values less than unity that decrease with increasing soil age (Alexander, 1974). Low Fe<sub>o</sub>/Fe<sub>d</sub> ratio indicates a high degree of crystallinity of Fe oxides (Obi et al., 2009). The nature, amount distribution of dithionite and citrate bicarbonate and acid ammonium oxalate extractable Fe and Al oxides in a soil significantly affect its properties, such as charge characteristics and ion adsorption, particularly phosphorus sorption (Obi et al., 2009). The fact that these oxides affect the physico-chemical properties of soils call for their study which would provide information that would be useful in planning, use and management of the soils, especially with respect to raphia palms which predominantly thrive in hydromorphic soils. Thus, the purpose of this study was to determine:

- the forms of Fe and Al oxides at the NIFOR Inland Valley earlier assessed for raphia palms cultivation and compare findings with an Upland soil in NIFOR.
- the interrelationship between Fe and Al oxides in both soils of the NIFOR Inland Valley and Upland.
- the interrelationship between Fe and Al oxides and some selected soil physicochemical properties of the Inland Valley and Upland soils of NIFOR.

# **MATERIALS AND METHODS**

# Field studies

Two locations were chosen; the NIFOR Inland Valley which is close to the Okhuo River located on Lat 6°34′N and longitude 5°43′E representing wetland or hydromorphic or poorly drained soil and the plantation behind chemistry Division, representing well-drained upland agricultural soil. Six (6) mini pits were dug on each location and two samples were collected from each pit, giving a total of twelve pits (12) per location. The mini pits were sunk in the various locations based on differences in vegetation, topography and surface visual soil colour. Thus variations in

the locations were adequately covered through purposive sampling. Samples obtained from the pits were stored in polythene bags and accurately labelled for further processing in the laboratory.

#### Laboratory study

The samples were taken to the laboratory and air dried. The air dried samples were crushed with the hand and sieved using a 2mm sieve. The samples that did not pass through this sieve were discarded while those that passed through were further analyzed for physicochemical properties as follows: Particle size analysis was determined using the hydrometer method using Calgon as a dispersing agent. Soil pH was determined in a 1:1soil/water suspension using a pH meter. Organic carbon was determined by the dichromate wet oxidation method of. The crystalline iron and aluminium oxides designated as Fed and Ald respectively were determined by dithionitecitrate buffered with sodium bicarbonate solution. The content of ammonium oxalate soluble iron and aluminium oxides also designated as Fe<sub>o</sub> and Al<sub>o</sub> respectively were determined by Mckeague and Day (1966) method.

#### Statistical analysis

Data were subjected to analysis of variance (ANOVA) using Genstat Statistical software. The means between depths were separated using Least Significant Difference (LSD) at 5% level of probability. The Matrix of Correlation Coefficient (r) was run using SPSS version 17.

#### RESULTS AND DISCUSSION

# Morphological Description of NIFOR Soils

The soils of NIFOR are inherently developed on Coastal Plain Sands. Soils formed on Coastal Plain Sands and Sandstones are similar in chemical characteristics, and have the same soil class Aguult in the USDA Soil Taxonomy. These soils are referred to as 'Acid Sands.' They are acidic, have low CEC, low base saturations and low fertility level, usually suffering from multiple nutrient deficiencies. Their prominent clay type is kaolinite. Four soil series of Acid Sands have been classified in NIFOR main station. They are Ahiara, Kulfo, Orlu and Alagba series. Soils of the Inland Valley are Ahiara series while those of the upland soils are Orlu series. Their detailed morphological descriptions are given in Table 1.

Table 1: Field Morphological Descriptions of Typical (Normal) NIFOR Soils

Profile No.	Horizon	Depth	Colour	Boundary	Structure	Texture***
And soil series		(cm)	(moist)	*	**	
5 Ahiara Series	A11	0-11	7,5YR3/4	a,w	1,c,g	Slightly clayey sand (LS)
(normal)	A12	11-36	5YR3/4	c,w	1,m,g	Clayey Sand (SL)
	B21	36-70	5YR3/3	d,w	1,m,g	Slightly clayey sand (LS)
	B22	70-104	5YR3/3	d,w	1,m,g	Slightly clayey sand (LS)
	B23	104-136	5YR3/3	d,w	1,m,g	Slightly clayey sand (LS)
	B24	136-174	2,5YR3/6	d,w	1,m,g	Clayey Sand (SL)
	B25	174-200	25YR3/6	d,w	1,m,g	Slightly clayey sand (LS)
10 Kulfo	A11	0-12	10YR3/2	c,w	1,c,g	Slightly clayey sand (LS)
Series(normal)	A12	12-40	7.5YR3/4	d,s	1,m,g	Slightly clayey sand (LS)
	A13	40-78	7.5YR3/4	d,s	1,m,g	Slightly clayey sand (LS)
	B21	78-105	5YR3/4	d,s	1,m,g	Clayey Sand (SL)
	B22	105-135	5YR4/6	d,s	1,m,g	Clayey Sand (SL)
	B23	135-163	5YR5/6	d,s	1,m,g	Clayey Sand (SL)
	B24	163-200	2.5YR5/6	-	2,m,sb	Clayey Sand (SL)
12 Orlu Series	A11	0-24	7.5YR3/4	c,s	1,m,g	Slightly clayey sand (LS)
(normal)	A12	24-55	2.5YR3/4	c,s	2,m,g	Clayey Sand (SL)
	B21t	55-88	2.5YR3/6	c,s	2,m,sb	Very clayey Sand (SCIL)
	B22t	88-126	2.5YR3/6	d,s	2,m,sb	Very clayey Sand (SCIL)
	B24t	126-154	2.5YR3/6	d,s	2,m,sb	Very clayey Sand (SCIL)
	B24t	154-200	2.5YR3/6	-	2,m,sb	Very clayey Sand (SCIL)
6 Alagba	A11	0-7	7.5YR3/4	c,s	1,m,g	Slightly clayey sand (LS)
Series	A12	7-21	5YR3/4	d,s	2,m,g	Clayey Sand (SL)
(normal)	B21t	21-40	2.5YR4/4	d,s	2,m,sb	Very clayey Sand (SCIL)
	B22t	40-67	2,5YR3/6	d,s	2,m,sb	Very clayey Sand (SCIL)
	B23t	67-105	2,5YR3/6	d,s	2,m,sb	Sandy clay (SCI)
	B24t	105-136	2,5YR3/6	d,s	1,c,sb	Sandy clay (SCI)
	B25t	136-200	2,5YR3/6	_	1,c,sb	Sandy clay (SCI)

<sup>\*</sup>a = abrupt, w = wavy, c = clear, d = diffuse, s = smooth

Source: NIFOR 1982

# The soil physico-chemical properties:

Some physical and chemical properties of the soils of both locations are shown in Tables 2 and 3. Soil pH and organic carbon at the Inland Valley increased with increased soil depth. In comparison with the upland soil, soil pH followed the same trend though it was slightly higher at the Upland. Soil pH was slightly acidic at the Inland Valley and moderately acidic at the Upland (Tables 2 and 3). Organic carbon was different in comparison with the Upland and decreased with increased soil depth. The increase in

organic carbon with increased soil depth at the Inland Valley could be due to the inefficient breakdown of organic matter by facultative anaerobes. This is consistent with the findings of Aghimien *et al.*, (1988) who called it lithological discontinuity. Comparatively, clay decreased with increased soil depth at the Inland Valley but increased significantly with increased soil depth at the Upland. Silt and sand followed the same trend in both locations.

<sup>\*\* 1 =</sup> weak, 2 = Moderate, c = coarse, m = medium, g = granular, sb = subangular blocky

<sup>\*\*\*</sup> LS = Loamysand, SL = Sandyloam, SCIL = Sandyclay loam, SCI = Sandyclay

# Distribution of dithionite and oxalate extractable forms of Fe and Al oxides in the soils

The citrate-bicarbonate dithionite extractable only indicate the Fe and Al that are present in the soil as free discrete bodies (i.e. crystalline and amorphous) but not those that are part of the structure of silicate minerals. These free discrete bodies are collectively referred to as uncombined or pedogenic free forms of the and could either be mobile element (amorphous) or immobile (crystalline). Tables 4 and 5 show the distribution of crystalline, definite crystalline and poorly crystalline form of Fe and Al oxides in soils of the Inland Valley and Upland. Fed ranged from 5.86 mg/kg to 151.38 mg/kg at the Inland Valley and increased with increased soil depth. In comparison with the Upland soil, it ranged from 11.13 mg/kg to 23.71 mg/kg and increased significantly with increased soil depth. Fed values were higher at the Inland Valley than the Upland soil values and were generally higher than Feo values of both locations. Fe<sub>o</sub> values of the Inland valley were lower than those of Upland which was probably due to the influence of water, a condition which inhibits weathering of

minerals. The increase in Fe<sub>d</sub> with increased soil depth was consistent with the findings of Juo and Maduakor (1974). They reported that dithionite Fe (total free Fe oxides) content increases with the increase of depth and added that the oxalate extractable Fe (amorphous Fe oxides) constitutes less than 10% of the total free Fe oxides throughout the profiles. Fe<sub>o</sub> ranged from 0.66 mg/kg to 9.64 mg/kg at the Valley and also increased with Inland increased soil depth. Fe<sub>o</sub> of Upland soil ranged from 5.01 mg/kg to 10.98 mg/kg and also increased with increased soil depth. Ald of the Inland Valley ranged from 8.20 mg/kg to 19.47 mg/kg and decreased with increased soil depth while Al<sub>o</sub> and Fe<sub>d</sub>-Fe<sub>o</sub> ranged from 0.32 mg/kg to 4.61 mg/kg and 2.34 mg/kg to 141.74 mg/kg respectively and both increased with increased soil depth. Al<sub>d</sub>-Al<sub>o</sub>, Fe<sub>o</sub>/Fe<sub>d</sub> and clay/Fe<sub>d</sub> ratios decreased with increased soil depth at the Inland Valley (Table 4). In Upland soil, Ald and Alo increased with increased soil depth and ranged from 7.28 mg/kg to 15.45 mg/kg and 2.64 mg/kg to 9.59 mg/kg respectively while the definite crystalline form decreased significantly with  $(Fe_d-Fe_o)$ increased soil depth. (Al<sub>d</sub>-Al<sub>o</sub>) decreased with increased soil depth.

Table 2: Physico-Chemical Properties of Nifor Inland Valley

S/No	Location/	рН (H2O)	O.C.	Clay	Silt	Sand	Textu
	Depth (cm) Inland V	(H2O)	(g/kg)	(g/kg)	(g/kg)	(g/kg)	ral Class
1	0-60	5.69	5.40	56.00	14.00	930.00	Sand
2	0-60	6.02	30.40	10.00	25.00	965.00	Sand
3	0-60	5.42	8.30	46.00	19.00	935.00	Sand
4	0-60	5.44	15.00	36.00	19.00	945.00	Sand
5	0-60	5.48	6.40	41.00	14.00	945.00	Sand
6	0-60	5.75	14.10	10.00	25.00	965.00	Sand
Mean	Mean	5.63	13.30	33.20	19.30	947.50	
7	60-120	6.31	8.00	36.00	24.00	940.00	Sand
8	60-120	5.39	25.30	20.00	35.00	945.00	Sand
9	60-120	5.85	3.50	41.00	14.00	945.00	Sand
10	60-120	5.79	35.50	20.00	25.00	955.00	Sand
11	60-120	5.48	6.40	41.00	14.00	945.00	Sand
12	60-120	5.35	18.90	20.00	30.00	950.00	Sand
	Mean	5.70	16.30	29.70	23.70	946.70	
	Range	5.35-6.31	3.50-35.50	10.0-56.0	14.0-35.0	930.0-965.0	
	LSD	NS	NS	NS	NS	NS	
	(0.05)						

NS = Non-significant.

Table 3: Physico-Chemical Properties of Upland Soils of Nifor

	Location/	pН	O.C.	Clay	Silt	Sand	Textural
S/No	Depth	(H2O)	(g/kg)	(g/kg)	(g/kg)	(g/kg)	Class
	(cm)						
	Upland						
1	0-60	5.87	8.0	36.0	19.0	945.0	Sand
2	0-60	6.12	7.0	36.0	14.0	950.0	Sand
3	0-60	5.99	8.0	41.0	24.0	935.0	Sand
4	0-60	6.18	8.0	41.0	29.0	930.0	Sand
5	0-60	6.28	8.6	36.0	29.0	935.0	Sand
6	0-60	6.35	7.4	31.0	9.0	960.0	Sand
	Mean	6.13	7.83	36.80	20.70	942.50	
7	0-60	6.27	6.1	56	24.0	920.0	Sand
8	0-60	6.21	5.1	86	24.0	890.0	Sand
9	0-60	6.19	5.4	76	19.0	905.0	Sand
10	0-60	6.2	5.1	61	24.0	915.0	Sand
11	0-60	6.32	5.4	86	24.0	890.0	Sand
12	0-60	6.37	4.5	36	24.0	940.0	Sand
	Mean	6.26	5.27	66.80	23.20	910.00	
	Ranges	5.87-6.37	4.5-8.6	31-86	14-29	945-890	
	LSD(0.05)	NS	NS	19.06	NS	17.79	

NS = Non-significant

Table 4: Dithionite and Ammonium Oxalate Iron and Aluminium Oxides of the Inland Valley of Nifor

S/No	Location/	Fed	Feo	Ald	Alo	Fed-Feo	Ald-Alo	Fe <sub>o</sub> /Fe <sub>d</sub>	Clay/Fe <sub>d</sub>
	Depth	(mg/kg)	(mg/kg)	(mg/kg)	(mg/kg)	(mg/kg)	(mg/kg)		
	(cm)								
1	0-60	10.01	7.67	19.47	0.55	2.34	18.92	0.77	0.56
2	0-60	111.57	0.66	13.25	2.42	110.91	10.83	0.01	0.01
3	0-60	10.5	3.92	11.05	1.21	6.58	9.84	0.37	0.44
4	0-60	27.49	0.8	8.97	1.76	26.69	7.21	0.03	0.13
5	0-60	6.66	1.17	15.12	2.86	5.49	12.26	0.18	0.62
6	0-60	39.93	5.69	9.19	5.6	34.24	3.59	0.14	0.03
	Mean	34.00	3.30	12.84	2.40	31.00	10.40	0.25	0.30
7	60-120	15.04	1.34	12.59	0.32	13.7	12.27	0.09	0.24
8	60-120	48.99	3.8	11.49	1.32	45.19	10.17	0.08	0.04
9	60-120	5.86	1.17	14.46	3.22	4.69	11.24	0.20	0.70
10	60-120	151.38	9.64	12.7	3.51	141.74	9.19	0.06	0.01
11	60-120	12.58	5.35	10.25	4.61	7.23	5.64	0.43	0.33
12	60-120	74.65	2.97	8.20	1.65	71.68	6.55	0.04	0.03
	Mean	51.00	4.00	11.61	2.44	47.00	9.20	0.15	0.23
	LSD(0.05)	NS	NS						

NS = Non-significant.

Table 5: Dithionite and Ammonium Oxalate Iron and Aluminium Oxides of Nifor Upland

	2011		-						
S/No	Location/	$Fe_d$	Feo	Al <sub>d</sub>	$Al_o$	Fe <sub>d</sub> -Fe <sub>o</sub>	Al <sub>d</sub> -Al <sub>o</sub>	Fe <sub>o</sub> /Fe <sub>d</sub>	Clay/Fe <sub>d</sub>
	Depth (cm)	(mg/kg)	(mg/kg)	(mg/kg)	(mg/kg)	(mg/kg)	(mg/kg)		-
1	0-60	23.71	5.72	15.45	5.82	17.99	9.63	0.24	0.15
2	0-60	18.45	5.01	7.28	3.18	13.44	4.1	0.27	0.20
3	0-60	20.05	7.12	11.57	3.62	12.93	7.95	0.36	0.21
4	0-60	34.46	7.47	9.59	4.61	26.99	4.98	0.22	0.12
5	0-60	19.19	7.04	7.47	5.49	12.15	1.98	0.37	0.19
6	0-60	20.39	8.98	10.92	3.18	11.41	7.74	0.44	0.15
	Mean	22.70	6.89	10.38	4.32	15.00	6.06	0.32	0.17
7	60-120	15.76	8.98	10.07	5.38	6.78	4.69	0.57	0.36
8	60-120	13.81	6.55	10.91	2.64	7.26	8.27	0.47	0.62
9	60-120	14.3	8.24	9.92	6.04	6.06	3.88	0.58	0.53
10	60-120	19.56	8.01	11.24	9.59	11.55	1.65	0.41	0.31
11	60-120	20.54	10.98	9.15	5.82	9.56	3.33	0.53	0.42
12	60-120	11.13	8.24	13.91	5.71	2.89	8.2	0.74	0.32
	Mean	15.80	8.50	10.87	5.86	7.30	5.00	0.55	0.43
	LSD(0.05)	5.65	NS	NS	NS	6.05	NS	0.13	0.12

NS = Non-significant.

Table 6: Correlation Matrix of Physico-Chemical Properties, Iron and Aluminium Oxides of inland valley soils of nifor

or iman	u vancy so	11 <u>8 01 1111</u>	<u>U</u> 1		<del>_</del>						
	Fed	Feo	Ald	$Al_o$	Fed-Feo	Al <sub>d</sub> - Al <sub>o</sub>	Fe <sub>0</sub> /Fe <sub>d</sub>	Clay/Fed	Oc	Clay	
Fed	-				<u> </u>						
Feo	0.340	-									
$Al_d$	-0.178	0.171	-								
$Al^o$	0.162	0.249	-0.309	-							
Fe <sub>d</sub> -Fe <sub>o</sub>	0.998**	0.284	-0.192	0.150	-						
$Al_d$ - $Al_o$	-0.208	0.035	0.922**	-0.653*	-0.214	-					
Fe <sub>o</sub> /Fe <sub>d</sub>	-0.511	0.448	0.592*	-0.132	-0.549	0.525	-				
Clay/Fe <sub>d</sub>	-0.71**	-4.153	0.628*	-0.124	-0.724**	0.551	0.617*	-			
Oc	0.945**	0.248	-0.252	0.075	0.948**	-0.231	-0.586*	-0.808*	-		
Clay	-0.720**	0.012	0.523	-0.389	-0.736**	0.575	0.750**	0.841**	-0.773**	-	

\*\*Correlation is significant at the 0.01 level, \*Correlation is significant at the 0.05 level, OC = Organic carbon.

Table 7: Correlation Matrix of Physico-Chemical Properties, Iron and Aluminium Oxides of Upland Soil in NIFOR

-	Fed	Feo	Ald	$Al_o$	Fed-Feo	Al <sub>d</sub> - Al <sub>o</sub>	Fe <sub>o</sub> /Fe <sub>d</sub>	Clay/Fe <sub>d</sub>	Oc	Clay
Fed	_				_	1110		=	-	
$Fe_{o}$	-0.097	-								
$Al_d$	-0.096	-0.065	-							
$Al^{o}$	-0.008	0.297	0.194	-						
Fe <sub>d</sub> -Fe <sub>o</sub>	0.968**	-0.345	-0.074	-0.083	-					
$Al_d$ - $Al_o$	-0.078	-0.263	0.734**	-0.523	-0.007	-				
Fe <sub>o</sub> /Fe <sub>d</sub>	-0.770**	0.607*	0.162	0.184	-0.880**	0.013	-			
Clay/Fed	-0.637*	0.278	-0.047	0.042	-0.671*	-0.070	0.614*	-		
Oc	0.640*	-0.412	-0.165	-0.329	0.708**	0.084	-0.769**	-0.774**	-	
Clay	-0.296	0.419	-0.159	0.169	-0.385	-0.252	0.375	0.891**	-0.647**	-

Correlation is significant at the 0.01 level, \*Correlation is significant at the 0.05 level, Oc = Organic carbon.

clay/Fe<sub>d</sub> ratios increased Fe<sub>0</sub>/Fe<sub>d</sub> and significantly with increased soil depth and ranged from 0.22 mg/kg to 0.74 mg/kg and 0.15 mg/kg to 0.62 mg/kg respectively. The Fe<sub>0</sub>/Fe<sub>d</sub> ratios were less than unity in both locations. It can be said that Fe<sub>d</sub> dominated the Inland Valley of NIFOR which is largely due to the migration of reduced components towards oxidized zones in the profile while Fe<sub>o</sub> dominated the Upland soils of NIFOR. The diagnostic index Fe<sub>0</sub>/Fe<sub>d</sub> clearly distinguished the poorly drained Inland Valley from the well drained upland soil. The fact that Fed dominated the Inland Valley soils of NIFOR is further buttressed by the Fe<sub>0</sub>/Fe<sub>d</sub> ratio which decreased with increased soil depth. Juo et al., (1974) reported a Fe<sub>o</sub>/Fe<sub>d</sub> ratio that decreased with profile depth and suggested that larger proportions of Fe oxides were present as crystalline forms in the lower horizons of the well drained soils that they studied. In both locations Fed was significantly correlated with Fe<sub>d</sub>-Fe<sub>o</sub> as was also observed by (Obi et al., 2009). The significant correlation between FED and Fe<sub>0</sub>/Fe<sub>d</sub> of both locations showed that most of the free fE oxides are present in crystalline forms. The significant correlation of Fed with organic carbon of both locations suggest that Fed was strongly associated with organic matter of both locations though the association was stronger at the Inland Valley as revealed in the high correlation value of 0.945\*\*. Feo had no correlation with Feo/Fed at the Inland Valley but was significantly correlated with Fe<sub>o</sub>/Fe<sub>d</sub> of Upland which again suggests the dominance of Fe<sub>o</sub> (amorphous Fe oxides) in Upland soils of NIFOR. Ald was significantly correlated with Ald-Alo of both locations which was consistent with the findings of Obi et al., (2009). This indicated that most of the free Al oxides are in the significantly crystalline forms  $Al_d$ was correlated with Fe<sub>o</sub>/Fe<sub>d</sub> and Clay/Fe<sub>d</sub> ratios of the Inland Valley but not with the Upland soils of NIFOR. There was no replacement of the crystalline Fe by Al as revealed by the lack of correlation between Fed and Ald of both locations. Like Fe<sub>d</sub> (total free Fe oxides), the definite crystalline Fe (Fe<sub>d</sub>-Fe<sub>o</sub>) also had significant correlations with organic carbon of

both locations. Fe<sub>o</sub>/Fe<sub>d</sub> was significantly correlated with clay at the Inland Valley (r = 0.705, P < 0.01) but was not correlated with clay of Upland soils. There was a lack of correlation between Fe<sub>o</sub>/Fe<sub>d</sub> and clay. There was no evidence of a co-migration of clay with Fe<sub>d</sub> in both locations though the negative correlation between Fe<sub>d</sub> and clay seemed to suggest a co-migration in Upland soils at the lower concentration of Fe<sub>d</sub>. The low mean values of Fe<sub>o</sub> of both locations indicate that the soils are highly weathered and contained very little weatherable minerals.

# Soils Supporting Raphia Palms

Raphia palms thrive best in deep soils, free of gravels or concretionary materials, of high moisture retention, preferably of shallow water table, coarse to medium texture and flat to gentle undulating topography to enhance less run-off and possible water logging (FMNAR, 1990). Alluvial and marine soils and the Acid Sands (Coastal plain sands, shale and sandstones) constitute the most important soils for raphia palm cultivation (FMANR, 1990; Aghimien 1982). These soils are widespread in Southern Nigeria where the palms play a major role in the traditional lives of the people. Their ability to support raphia palms stem from the fact that these soils are low in sesquioxide contents. Aghimien (1982) in his study of hydromorphic soils supporting raphia palms in Southern Nigeria reported that the low amount of P in occluded forms was partly attributed to the low contents of sesquioxides which were responsible for occluding P fractions under water logged and reduced conditions. This present study has revealed that soils of NIFOR can adequately support raphia palms cultivation as the dithionite and oxalate extractable sesquioxides contents are low and may not sufficiently occlude those nutrients needed by raphia palms in high amounts.

#### **CONCLUSION**

Crystalline forms as compared to the amorphous form dominate the free sesquioxides in soils of NIFOR main station as Fe<sub>d</sub>, Fe<sub>d</sub>-Fe<sub>o</sub>, Al<sub>d</sub> and Al<sub>d</sub>-Al<sub>o</sub> contained

38.91%, 35.57%, 11.90% and 8.90% of the total sesquioxide contents of NIFOR Inland Valley respectively. Feo and Alo contained 3.34% and 2.20% of the total sesquioxide content respectively. This trend continued with the Upland soils. These amounts are however not high enough to cause a distortion of the physical property such as hardness of the soils but high enough to influence chemical properties of the soils such as ion adsorption and aggregate stability. Furthermore, the high significance of these crystalline forms of Fe with organic carbon in both locations suggests that organic matter can help mitigate their effects on the physical and chemical properties of the soils. Also significant is that Fe<sub>o</sub>/Fe<sub>d</sub> ratio clearly distinguished the well drained Upland soil from the poorly drained Inland Valley. There was no evidence of co-migration of clay with Fed as revealed by the lack of correlation between Fed and clay of both locations. Further studies are needed to study the association of these oxides with nutrients such as N, P, K, Mg and Zn which are highly needed by raphia palms.

# **REFERENCES:**

- Aghimien A.E. (1982). The Chemistry and Fertility Status of Hydromorphic Soils Supporting Raphia Palms (Raphia spp) in Southern Nigeria. Ph.D Thesis, University of Ibadan, Ibadan Nigeria.
- Aghimien A.E., Udo, E. J. and Ataga D.O. (1988). Profile distribution of forms of Iron and Aluminium in the Hydromorphic soils of Southern Nigeria. *Journal of West Africa Science Association 31*: 57-70.
- FMANR (1990). Literature Review on Soil Fertility Investigations in Nigeria (in Five Volumes. Bobma Publishers, U.I. P.O. Box 9555, Ibadan.
- Juo A.S.R. and H.O. Maduakor (1974). Forms of pedogenetic distribution of extractable iron and aluminium in

- selected soils of Nigeria. *Geoderma* 11:167-179
- Mckeague, J.A. and Day, J.A. (1966).

  Ammonium oxalate and DCB extraction of Fe and Al. Canada Journal of Soil Science. 46: 13-22
- McKeague, J.A.; Nowland J.E.; Brydon J.E. and N.M. Miles (1971). Characterization and classification of five soils from eastern Canada having prominently mottled B horizons. Can J. Soil Sci. 51: 483-497
- Mehra O.P. and M.L. Jackson (1960). Iron oxide removal from soils and clays by dithionite-citrate system buffered with sodium bicarbonate. Proc. 7th Natl. Conf. Clays Clays Miner. Pergamon Press, New York, pp. 317-327.
- Obi, J.C.; Akinbola, G.E. and Anozie H.F. (2009). Distribution of dithionite and oxalate-extractable iron oxides of a catena in the basement complex soils of Southwestern Nigeria. *Nigerian Journal of Soil Science*19
- Ogunsola O.A.; Omueti J. A.; Olade O. and Udo E.J. (1989). Free Oxide status and distribution in soils overlying Limestone areas in Nigeria. *Soil Science* Vol. 147, No 4
- Osodeke, V.E.; I.L. Nwotiti, B.O. Nuga (2005). Sesquioxides Distribution along a Toposequence in Umudike Area of Southeastern Nigeria. Electronic Journalof Environmental, Agricultiral and Food Chemistry 4(6)[1117-1124].
- Schwertmann, U. (1964). Differenzierung der Eisenoxyde des Bodens durch Extraktion mit Ammoniumoxalat-Losung. Z. Pflanzenernahr. Bodenkd., 105: 194-202.